

Electrical Resistivity Investigation for Ground Water in Parts of Pegi, Federal Capital Territory, Nigeria

EnahoroIfidon Asije¹, Ogbonnaya Igwe²

¹Department of Geology, University of Nigeria, Nsukka, Enugu State, Nigeria

²Department of Geology, University of Nigeria, Nsukka, Enugu State, Nigeria

Abstract: Geophysical investigation involving the use of vertical electrical sounding (VES) technique was employed in the search for potential ground water target at twenty (20) locations within parts of Pegi, in Kuje Area Council of the Federal Capital Territory, Nigeria. The results of the survey indicated the presence of at least two layers of overburden (1.2-32m thick and $10\Omega\text{-m}$ to $2,300\Omega\text{-m}$) before the bedrock. The resistivity of the basement varies from $168\Omega\text{-m}$ to $6,300\Omega\text{-m}$. A transition zone of considerable thickness composed majorly of partly weathered bedrock underlies about two-third of the surveyed area. The two major aquifers in basement terrain i.e. weathered (overburden) and fractured bedrock were delineated from the geophysical investigation. Boreholes drilled within this area intercepted both the weathered and fractured bedrock aquifers. The potential for ground water in this area is good and the yield is sufficient enough to meet the demands of the clients.

Keywords: Aquifers, Resistivity, Transition zone, Vertical Electrical Sounding (VES), yield

I. Introduction

From its creation in 1976 and up until the early 1990's when the final relocation of the nation's capital from Lagos was effected the Federal Capital Territory, F.C.T. (Fig. 1)) has witnessed an unprecedented influx of people who have besieged the new seat of power for job opportunities both with the federal government and private establishment.

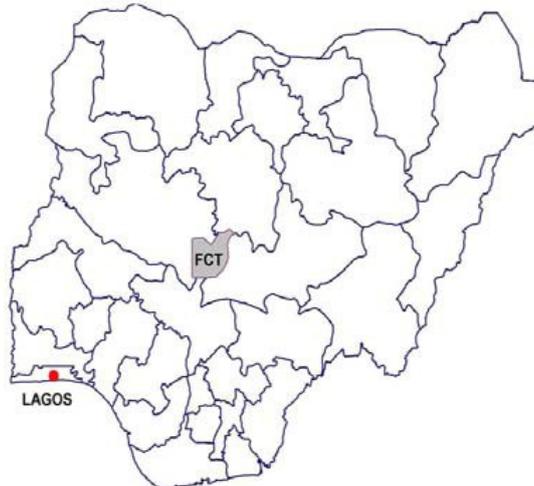


Figure 1: Map of Nigeria Showing The Location Of Lagos And The New Federal Capital Territory, (AGIS, 2006) (Culled From Jibril, 2009)

Also included in this mass movement of people are professionals and artisans who are independently offering services to the teeming population of the F.C.T. The astronomical rise in the population of the F.C.T. has led to massive shortage of basic infrastructure such as water supply. The inability of the existing Usman dam to adequately provide portable waterto meet the demand of those in the city center and environs has led the inhabitants of the FCT to source for portable drinking water at their own expense. One of such ways is through the drilling of water boreholes. In order to reduce the risk of drilling abortive wells electrical resistivity survey was therefore employed in the search for ground water in parts of Pegi, Kuje Area Council, F.C.T. (Fig. 2).



Figure2: Satellite map showing parts of Pegi (from Google Earth)

The electrical resistivity investigations discussed in this paper were carried out over a period of about two (2) years spanning between December, 2012 and May, 2014 with a total of twenty (20) survey stations (Table 1). The essence of this paper is to highlight the success of the electrical resistivity survey in delineating both weathered and fractured bedrock aquifers which have been successfully drilled and put to use in the said locality. And also make this information available to hydrogeoscientists who may want to carry out similar prospecting or drilling within the said location. It should be noted that the authors were not responsible for the drilling as they were only contracted by borehole contractors to carry out the electrical resistivity investigation. However, information about the drilling process was verbally obtained from the drillers.

II. Geology And Hydrogeology

The area under consideration falls within the basement complex of north-central Nigeria (Fig.3).

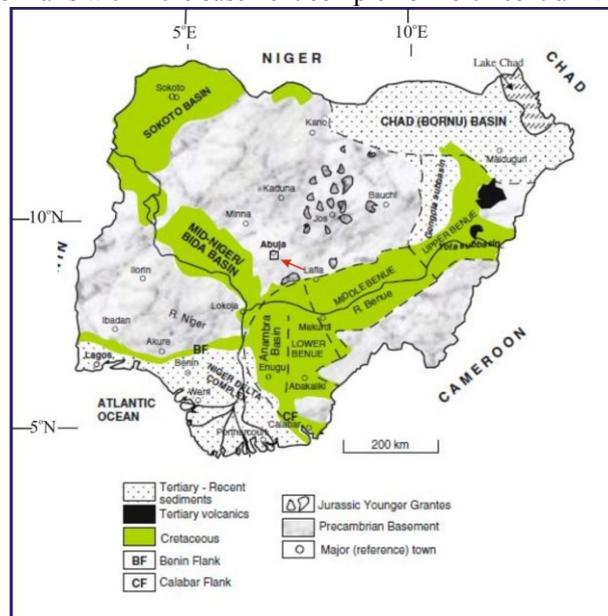


Figure 3: Geological Map of Nigeria, showing the position of Abuja (red arrow) in the Basement Complex of North central Nigeria (modified from Obaje, 2009). (cullled from Omeje M. et al, 2013)

The rocks include different textures of granites, coarse to fine, consisting essentially of biotite, feldspars and quartz (Eduvie etal, 2003). In most cases the rocks have weathered into reddish micaceous sandy clay to clay materials capped by laterites (Offodile 1992).The main sources of water in basement complex are the weathered/transition zone and fractured basement rocks.

III. Materials And Method

The popular VERTICAL ELECTRICAL SOUNDING (VES) of electrical resistivity method was employed. Basically, the method was designed to detect the physio-structural characteristics of basement rocks such as variation in conductivity and thereby determine the depth of occurrence of productive aquifers. The survey involved the selection of 20 geo-stations and probing points. Schlumberger array was employed using Allied Ohmega Resistivity Meter. Coordinates of VES points were obtained using Garmin's Geographical Positioning System (GPS) Personal Navigation unit. The field data obtained were interpreted using IPI2Win Resistivity Sounding Interpretation Software (1990-2008). Surfer 10 from Golden Software, Inc. was used to enhance the presentation of the interpreted data.

IV. Results And Discussion

The data set analysis of all the sounded points indicated at least two layers of regolith or overburden before the bedrock. Typical curve types analysed from the field data are KH, H and HQ. The KH and H-type curve accounted for 85% of the total curve types (Table 1 and Fig.4a-c).

Table 1: Summary of VES Interpretation

Date	VES No.	UTM		Layer resistivity				Depth			Inferred Lithology	Curve Type	Overburden thickness
		Easting	Northing	ρ_1	ρ_2	ρ_3	ρ_4	d1	d2	d3			
24/12/12	1	311186	975861	1424	151	359	-	1.08	11.7	-	hardpan, sandy clay and fractured basement	KH	11.7
12/01/13	2	310506	976359	822	172	1271	-	0.6	10.1	-	laterite, sandy clay, fractured basement	H	10.1
4/02/13	3	310402	976351	627	220	81	719	0.5	7.5	14	clayey sand, sandy clay, fractured basement	H	13.9
6/03/13	4	310567	976100	252	102	39	2214	1.2	7.2	14	sandy clay, clay, fractured basement	H	13.6
11/03/13	5	310354	976305	531	132	718	168	1.2	3.4	14	sandy clay, clayey sand, fractured basement	HQ	3.5
21/03/13	6	310520	976355	368	109	33	3366	1.7	6.2	32	sandy clay, clay, fairly fractured basement	H	32.3
11/04/13	7	310617	976009	202	415	21	245	0.9	1.7	5.2	sandy clay, clayey sand, clay, fractured basement	KH	5.2
17/06/13	8	310628	976026	130	518	9.1	1248	1.2	2.9	7.5	sandy clay, clayey sand, clay, fractured basement	KH	7.5
17/08/13	9	309742	976427	107	462	160	6433	0.5	10.7	21	Sandy clay, clayey sand, fractured basement	KH	20.7
26/08/13	10	310672	976283	57	1583	28	706	0.5	1.1	5	clay, laterite, fractured basement	KH	5
02/10/13	11	310767	975697	350	74	2150	1850	0.5	0.6	5.5	sandy clay, clay, fractured basement	HQ	1.2
15/02/14	12	310527	976467	225	1025	-	-	14.2	-	-	sandy clay, fractured basement	A	14.2
15/03/14	13	310229	976449	1517	168	93	1286	0.7	2.5	17	laterite, sandy clay, fractured basement	H	17.3
19/03/14	14	310250	975865	706	2207	206	1682	0.5	1	14	clayey sand, laterite, sandy clay, fractured basement	KH	14.2
27/03/14	15	309972	975885	2348	10	1534	-	0.9	7.3	-	laterite, clay, fractured basement	H	7.3
08/04/14	16	309998	975991	328	266	1426	-	0.5	12.4	-	sandy clay, fractured basement	H	12.4
08/04/14	17	310428	976461	453	266	16	1544	0.7	5.2	12	clayey sand, clay, fractured basement	KH	12.4
19/04/14	18	309907	975945	275	409	124	907	0.5	3.6	9.6	sandy clay, clayey sand, fractured basement	KH	9.6
19/04/14	19	309824	976028	116	327	192	725	0.5	1.6	5.2	sandy clay, fractured basement	KH	5.2
12/05/14	20	309964	975920	66	229	51	399	0.5	4	9	sandy clay, clay, fractured basement	KH	13.6

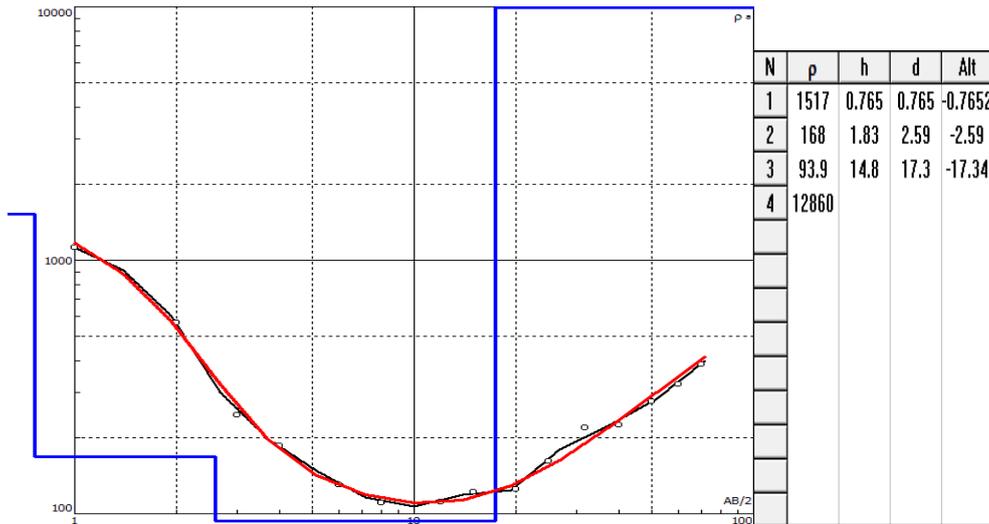


Figure 4a: H-type curve

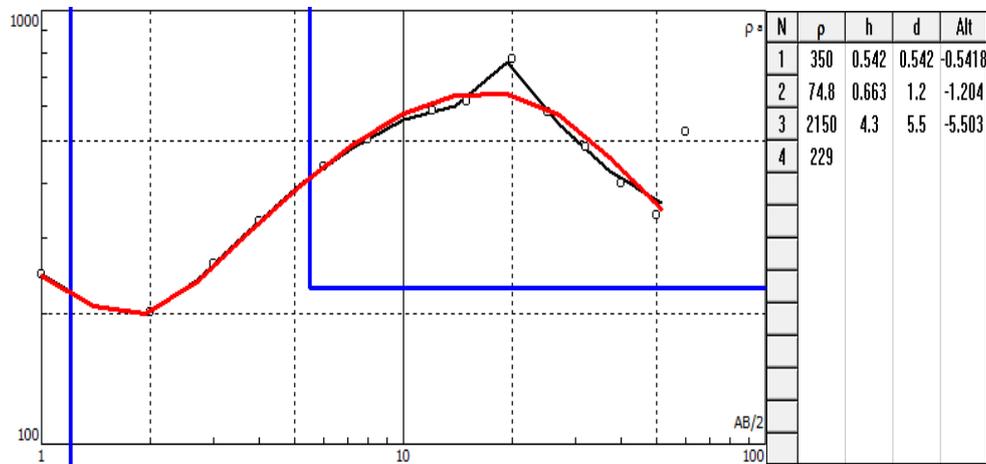


Figure 4b: HQ-type curve

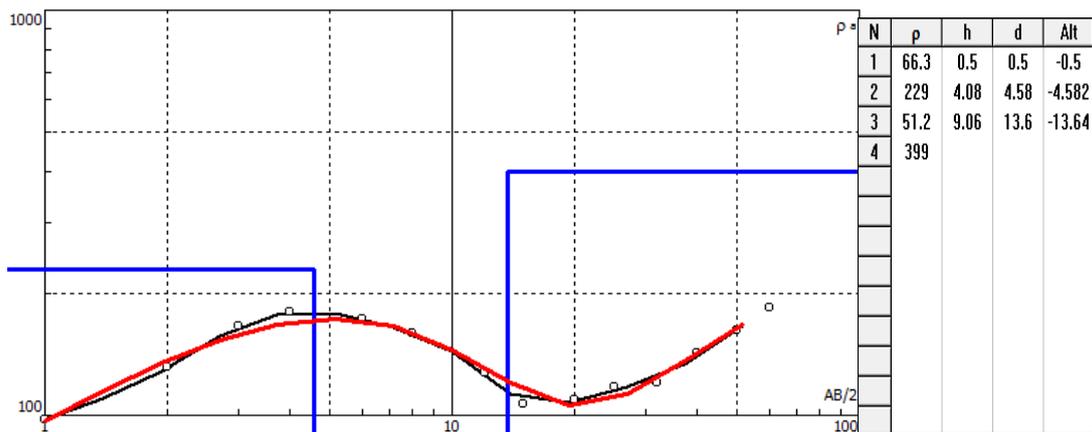


Figure 4c: KH-type curve

4.1 ISO-Pach And Iso-Resistivity Map Of The Overburden

Overburden thickness viz-a-viz depth to bedrock varies from 1.2 meters at the shallowest to about 32 meters Fig.5). Overburden is thickest at the north-central end of the area marked "T" and shallowest at the south-

eastern end of the area marked "S" where the overburden thickness is almost negligible.

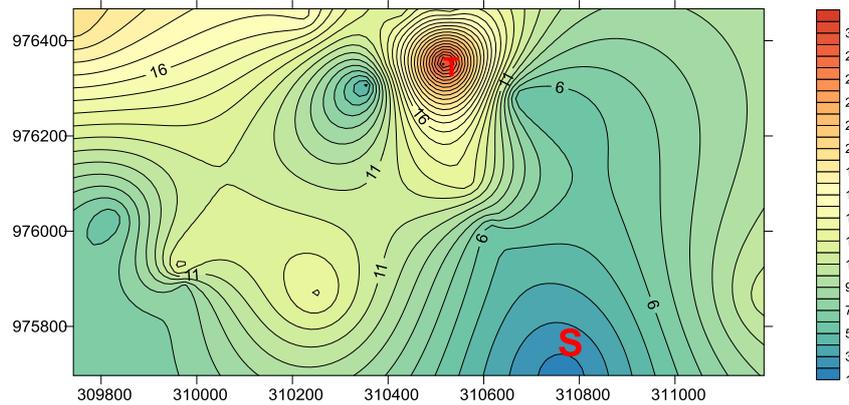


Figure 5: Iso-pach map of the overburden

Average overburden thickness around the area is about 10 meters. The area marked "T" also infer a depression or trough within the bedrock configuration with the properties of a reservoir capable of storing underground water. And as such the thick overburden at this area would act as good aquifer for underground water development.

Resistivity of the overburden varies from as low as 10Ω-m at the north-central part of the area to as high as 2,300Ω-m (Fig.6 and Table 1).

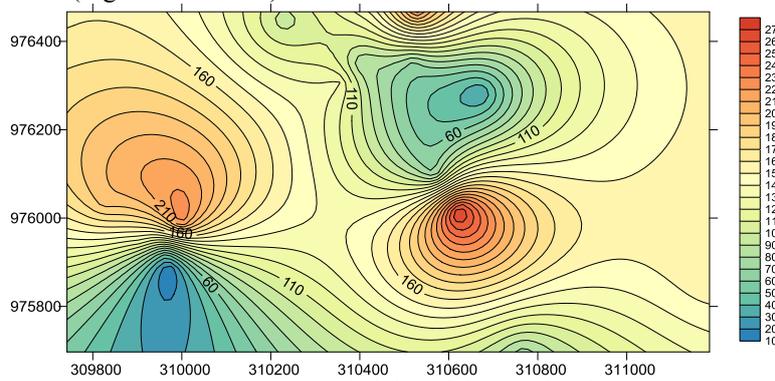


Figure 6: Iso resistivity of the overburden

Inferred lithologic units of the overburden delineated on the basis of resistivity values obtained are laterite, hardpan, sandy clay, clay and clayey sand (Table 1). Wherever the laterite and hardpan were delineated at any given VES point they usually appear within the first two layers i.e. from the surface.

1.2 ISO-Resistivity Of The Bedrock

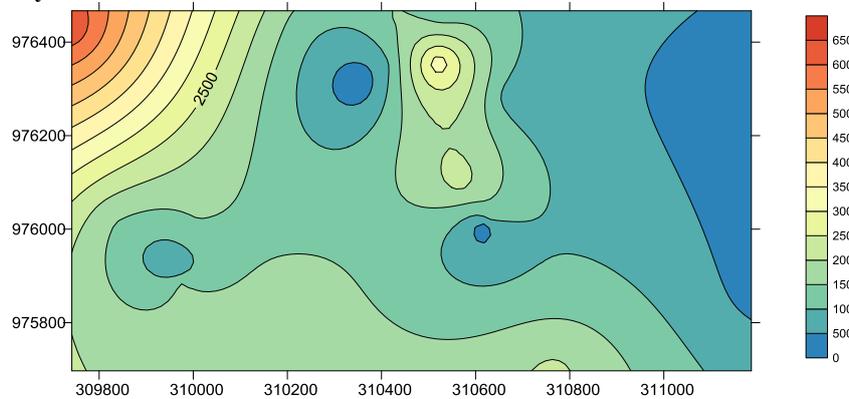


Figure 7: Iso resistivity of the basement

Resistivity of the bedrock varies from 168Ω-m to 6,300Ω-m (Fig.7 and Table 1).The bedrock resistivity depicts a geological formation with well-developed secondary fractures i.e. bedrock aquifers, across the entire surveyed points. Depth to this zone varies from as low as 2 meters to as high as 32 meters (Fig.5 and

Table 1). A transition zone composed of partly weathered bedrocks between the overburden and bedrock was inferred from the iso-resistivity map on areas with bedrock resistivity of less than $1500\Omega\text{-m}$. This zone appears to cover more than two-third of the entire survey area. And as such, no distinct contact between the overburden and bedrock appear to exist within such locations. It was therefore recommended that wells to be drilled within this zone should be cased to the bottom and well gravel-packed to forestall future collapse or caving.

Hard, competent and fractured bedrock is expected to be encountered at the extreme north-western part of the area (Fig.7).

The bedrock therefore has good potential for underground water development. Information obtained from the drilling contractor corroborated the findings above. All the wells drilled have since been commissioned and put to use by the individual clients. Yields were satisfactory.

V. Conclusion

Electrical Resistivity investigation for ground water was carried out at some parts of Pegi in Kuje Area Council of the F.C.T. Nigeria. The method proved successful in delineating both weathered and fractured bedrock in relation to aquifer potential of the area and guided the location of drill points for productive water boreholes.

It was however observed that where ever the overburden had a considerable thickness i.e. transition zone, the contact between the overburden and the bedrock was difficult to delineate.

Hydrogeoscientists who may want to carry out similar prospecting or drilling program within the said locality should put this into consideration while making recommendations for drilling by ensuring that provision is made for down-to-bottom casing and gravel-packing to forestall future collapse or caving of the drilled well.

References

- [1]. Ibrahim UsmanJibril (2009). Squatter Resettlement/ Relocation Programme in Abuja, Nigeria and the Issue of Land Title Security. FIG Working Week 2009, Surveyors Key Role in Accelerated Development Eilat, Israel, 3-8 May 2009
- [2]. 2013 Google Earth Inc.
- [3]. Omeje M., Husin W., Noorddin I., Oha I. A., Onwuka O.S, Ugwuoke P. E. and Meludu (2013). Geoelectrical investigation of aquifer problems in Gosa area of Abuja, North Central, Nigeria. Int. J. Phys. Sci. Vol. 8(13), pp. 549-559
- [4]. [M.O. Eduvie, T. Olabode and O.O. Yaya, Nigeria (2003). Assessment of groundwater potentials of Abuja environs. Towards the Millennium Development Goals, 29th WEDC International Conference Abuja, Nigeria. pp 183-187
- [5]. M.E. Offodile, An Approach to Groundwater Study and Development in Nigeria (Mecon Services Ltd., Jos, 1992)